

An Assessment of Organic Carbon Fractions in Paddy and Associated Non-paddy Soils of Upper Brahmaputra Valley of Assam

Zenesia A. Phillips* and R. M. Karmakar

Department of Soil Science, Assam Agricultural University
Jorhat, India

*Corresponding author's email: zenesia_phillips [AT] yahoo.com

ABSTRACT----- An investigation was carried out to study the soil organic carbon fractions in paddy and associated non-paddy soils of Assam, India. Three districts viz. Sivasagar, Jorhat and Golaghat in the Brahmaputra valley of Assam were selected for this study. Soil samples were collected from six profiles, three each from mono-cropped paddy and associated non-paddy areas were collected. Horizon-wise, soil samples were analyzed for organic carbon fractions. Organic carbon, Humus C, Humin C, Humic acid C and Fulvic acid C were all analyzed from the samples. Organic carbon content in soil varied from 0.90 -7.90g kg⁻¹. Surface horizons of paddy soils contained lower amounts of humus C (0.255-0.330 per cent) and fulvic acid C (C_{FA}) (0.135 – 0.180 per cent) and higher amounts of humin carbon (0.210 - 0.475 per cent) and humic acid C (C_{HA}) (0.120-0.150 per cent) as compared to that of non-paddy soils (0.300-0.435 and 0.195-0.300, 0.180-0.490 and 0.105-0.135 per cent respectively). It was concluded that the ratio of (C_{HA}+C_{FA})/C_{TOC} indicated less humification on the surface horizons of paddy soils. This is a direct result of prolonged submergence which led to humic acid carbon becoming less oxidized or humified.

Keywords--- Organic carbon, humic acid, fulvic acid, humin, humification, paddy soils

1. INTRODUCTION

Soil organic carbon (SOC) is a heterogeneous mixture of simple and complex organic carbon compounds which comes from the products of living organisms. It is the carbon stored within soil and is a part of soil organic matter which comprises of other important elements such as calcium, hydrogen, oxygen and nitrogen. There are four biologically significant types of soil organic carbon fractions identified, each differing in size, hence causing decomposition to go on at different rates. These fractions are: crop residues, particulate organic carbon, humus and recalcitrant organic carbon (Baldock, 2008).

Soil organic carbon plays a vital role in improving soil quality for sustainable crop production. Studies relating to humic substances in soil are important for understanding both soil genesis and management of soil organic matter (SOM). Soil organic carbon is important as it determines ecosystem and agro ecosystem functions, influencing pedogenesis. Greater amounts of organic carbon in paddy soil can be a reflection of intensified humification resulting from increased aggregation (Wang *et al.*, 2015).

It is also of global importance because of its role in the global carbon cycle and, therefore, plays a role in the mitigation of atmospheric levels of green house gases (GHG) with special reference to carbondioxide (CO₂) due to global warming (Bhattacharya *et al.*, 2008). Organic matter (OM) in soil represents the largest carbon pool and is an important nutrient reservoir in terrestrial ecosystems. Nearly 70 per cent area of India is deficient in soil organic carbon (Velayutham *et al.*, 2000).

Rice (*Oryza sativa* L.), grown in India is a major cereal crop of Asian origin. Rice occupies about two thirds of the total cropped area in Assam. It is the single major food source, hence significantly contributes to mass agricultural production, therefore playing a significant role in the state economy. Rice in Assam is grown as autumn (*ahu*), winter (*sali*) and summer (*boro*) crop based on the combination of land and hydrological characteristics, maturity duration of rice genotypes, length of growing season and growing conditions. Out of these, *sali* rice grown during the wet season (June-December) is the most important culture occupying about 70 per cent of the total rice area of state of Assam. This major portion of rice cultivated area particularly in the Upper Brahmaputra Valley Zone (UBVZ) of Assam is under mono-crop (Barah, 2001). Paddy soils are a unique anthropogenic soil type formed under long-term hydro-agric management with seasonal submergence (Gong, 1999). The submergence and puddling during rice cultivation creates conditions which are significantly different from those of non-rice growing areas (Willet, 1979; Gong, 1983).

Undoubtedly, rice cropping, as an important land use, has significant effects on C, and N cycling around the globe (Lal *et al.*, 1998). It is therefore imperative that studies be carried out in relation to how mono-cropping affects soil organic carbon fractions. Studies of this nature are necessary in assessing damage done to soil over time, and guide in finding ways to

improve the soil fertility, leading to soil sustainability and continued productivity of this crop. It is commonly accepted that implementation of rice-cropping practices would enhance accumulation of soil organic carbon (Lal, 2002)

Soil organic matter (SOM) is a complex mixture of plant and animal residues under various stages of decomposition. These residues or substances are synthesized microbiologically and/or chemically by microorganisms. The organic materials are colonized by microbes which utilizes enzymes to oxidize it, while obtaining energy and carbon dioxide. (Schnitzer and Khan, 1972). The composition of organic matter is extremely complex because of the nature of the various organic inputs and their different stages of decomposition (Chenu *et al.*, 2014).

Paddy soils are an important accumulator for organic matter (Zhang and He, 2004). The accumulation of organic carbon in soils is influenced by the type of plant material. Sanchez *et al.* (1982) observed that there was higher organic matter content in tropical soil under forest vegetation as compared to grassland vegetation. However, it was found that the organic matter content under grassland vegetation was higher in temperate regions. Cultivation adversely affects the organic matter content in soil. Soils under normal agricultural practices are subject to be depleted in their organic matter content; however this greatly depends on the management practices carried out in the field (Rudrapa, 1978). Borah and Karmakar (1999) observed that the soil organic carbon content under plantation crop was higher than that under the agricultural crops which might have been attributed to annual addition of more amounts of leaf litters in the soils under plantation crops. But the cultivated soils were characterized by lower quantity of organic matter status as compared to forest and grass vegetation (Karmakar, 1999). Management of paddy soil is believed to be favorable for accumulation of organic matter and its content in paddy soils were found to be statistically higher than that of non-paddy soils (Cai, 1996). In Southeast China, paddy soils have the second largest organic matter stocks and thus a large proportion of terrestrial carbon is conserved in wetland rice soils (Zhao, 1996). It is therefore believed that implementation of rice-cropping practices would enhance accumulation of soil organic C (Lal, 2002).

Organic materials decreased exponentially with increasing depth in paddy soils (Pan *et al.*, 2008).

Soil organic carbon (SOC) decreased exponentially with depth to 100 cm in paddy soils; a substantial proportion of the total SOC (30-40%) was stored below the 30 cm depth. This result suggests that SOC stratification within profiles varies with different pedogenetical types of paddy soils with regards to clay and iron oxyhydrates distribution (Genxing *et al.*, 2008). Soil organic matter accumulation in paddy sub-soil can be explained by downward movement of dissolved organic matter and its stabilization by interaction with iron oxides (Kogel-Knabner *et al.*, 2010). Soil organic carbon (SOC) stratification within profiles varied with different pedogenetical types of paddy soils with regards to clay and iron oxyhydrates distributions, therefore organic materials decreased exponentially with increasing in depth in paddy soils (Genxing *et al.*, 2008).

Soil organic matter can be divided into several fractions depending on their densities. Labile fraction (LF) is the most prominent, partly due to its high turnover rate plus it is easily affected by management systems as well as erosion (Wong *et al.*, 2014).

Labile fractions has been described in various ways by soil scientists, including particulate organic carbon (POC) (53–2000 μm), light fraction organic carbon (LFOC) (density of $< 2.0 \text{ g cm}^{-2}$), readily oxidized carbon (ROC) (easily oxidized by potassium permanganate), soil microbial biomass carbon (SMBC) and dissolved organic carbon (DOC) (Cao *et al.*, 2013).

Schnitzer and Khan (1972), grouped soil organic matter (SOM) into two groups - humic and non humic substances. Humic substances are the dark coloured, amorphous macromolecules, ranging in molecular weight from a few hundred to several thousand; they make up the major organic fraction in soil, which are synthesized by biological, chemical and physical processes. Hayes and Swift (1978), found that humic substances constitute 50 to 80 per cent of SOM and are considered to include the most stable fractions therefore, common fractionations of SOM are based on differences in solubilities of organic constituents under acid and alkali conditions. Humic substances represent approximately 40-60% of the soil organic matter, including three different fractions, which are defined according to their different stabilities under acid hydrolysis and permanganate oxidation (Paul *et al.*, 2001). Humic acid (HA) is said to be the insoluble fraction of humic substances; humic acid (HA) is the fraction that is soluble under alkaline conditions; and fulvic acid (FA) is the fraction that is soluble under both alkaline and acidic conditions (Sutton and Sposito, 2005). The acid soluble material classified as fulvic acid, invariably contains organic substances classified as non-humic (Stevenson and Elliott, 1989). These three humic fractions are of similar structure but differ in molecular weight and functional group content. Humic substances are more stable organic compounds that make up a significant portion of the total organic C and N in soil (Lal, 1994). Specific fractions of organic matter play an important role in maintaining soil quality and thus could be important for indicating and assessing the impact of management practices (Cambardella and Elliot 1992; Chan, 1997).

Milori *et al.* (2002) reported an increase in the percentage of humic substances with soil depth across all land uses. Long-term rice cropping has been proven to increase organic C content and humic acid/fulvic acid ratio; it also causes downward movement of organic C, N, and P, which may result in a number of environmental impacts (Zhang and He, 2004). Studies show that the distribution of humic substances in the soils developed on different physiographic units in lower Brahmaputra valley zone of Assam. Fulvic acid majorly constituted the largest part (47.6-90.0%) of humus carbon. Humic acid content decreased gradually with soil depth whereas fulvic acid was found to increase in the illuvial

horizons of Ruptic-Ultic Dystrudepts on piedmont plain and Typic Hapludults on monadnocks (Karmakar and Rao 1999).

2. MATERIALS AND METHODS

Materials:

Location and extent: Three districts viz., Sivasagar, Jorhat and Golaghat in the Brahmaputra valley of Assam were selected for the present study. These districts are situated in the Upper Brahmaputra Valley Zone (UBVZ) of Assam and cover an area of 9021 sq. km which is 55.7 per cent of UBVZ. These districts form a part of the southern bank of the Brahmaputra valley of Assam. Six soil profiles, three each from mono-cropped paddy and associated non-paddy areas were collected from Sivasagar, Jorhat and Golaghat districts of Assam. Horizon-wise soil samples were collected from each soil profile.

Preparation of soil samples and analysis: The soil samples were air dried, ground and passed through a 2 mm sieve. The sieved soil samples were stored in polythene bags and subsequently used for various physico-chemical analyses. Fresh soil samples were stored in refrigerator for microbiological analyses.

Organic carbon was determined by the Walkley and Black's method Jackson (1973)

Sub-fractionation of organic carbon was done by Modified Walkley and Black Chan *et al.* (2001) method.

Humus (CH), humin (CHn), humic acid carbon (CHA) and fulvic acid carbon (CFA) fractions: Humus carbon content in the 0.1 M Na-pyrophosphate extract was determined by Walkley and Black (1934) method. The difference between organic carbon and humus carbon gives the humin carbon. The extract was acidified with conc. H₂SO₄ to precipitate humic acid fraction (Kononova *et al.*, 1966); the precipitate was filtered and dissolved in warm 0.05 N NaOH solution. Carbon content in the acid soluble (fulvic acid fraction) and alkali soluble (humic acid fraction) were determined as earlier (Kononova *et al.*, 1966).

Statistical analysis

Simple correlation analyses were carried out for some selected soil parameters following the procedure outlined by Snedecor and Cochran (1967).

Parameters of interest included soil organic carbon fractions:

Organic Carbon (OC%)

Humus Carbon (C_H)

Humin C_{Hn})

Humic acid (C_{HA}) and

Fulvic acid (C_{FA})

3. RESULTS AND DISCUSSION

Soil organic carbon fractions: The data on soil organic carbon fractions are presented in Table 1.1. and Table 1.2. The amount of humus carbon (C_H) in soil varied from 0.045-0.330 per cent in the paddy soils (P1, P2, P3) and 0.030-0.435 per cent in the non-paddy soils (NP1, NP2, NP3) (Table 4.8). Surface horizons of paddy soils contained lower amount of humus carbon (0.255-0.330 per cent) as compared to that of non-paddy soils (0.300-0.435 per cent). Humus carbon (C_H) constituted 16.4 -61.1 per cent of soil organic carbon (SOC) in the paddy soils and 38.0-69.0 of soil organic carbon (SOC) in non-paddy soils (Table 1.2.).

This suggests lower degree of humification in the surface horizon of paddy soils under aquatic condition. This is also supported by higher amount of humin C and lower values of (C_{HA}+C_{FA})/C_{TOC} ratio in the surface horizons of paddy soils as compared to that in the surface horizons of non-paddy soils (Table 1.1.). The humus C was significantly and positively correlated with soil organic C (r= 0.868**) and exchangeable Ca⁺⁺ (r= 0.508**) suggest that the degree of humus formation depends on the amount of soil organic carbon and exchangeable Ca⁺⁺. Soil acidification results in a significant reduction of decomposition rates of organic matter and increases soil carbon accumulation. (Paustian *et al.*, 1997). Research has shown a 20% increase in soil exchangeable Ca⁺⁺ due to increased organic fertilizer application (Yan and Hou, 2018). It can therefore be suggested that rice cultivation done over time with the implementation of organic related practices played a major role in increasing SOM while increasing exchangeable Ca⁺⁺. More exchangeable Ca helps to alter soil acidity, since the exchangeable Ca⁺⁺ ions are attached to the cation exchange sites of the organic matter particles and is strongly absorbed in the soil complex. On the other hand, soil porosity and soil acidity retard the process of humification as indicated by negative correlation of humus C with bulk density (r= - 0.607**) and pH (r= - 0.341) (Table 1.3.).

The amount of humin carbon (C_{Hn}) in soil varied from 0.011-0.475 per cent in the paddy soils and 0.030-0.490 per cent in the non-paddy soils. Surface horizons of paddy soils contained higher amount of humin C as compared to that of non-paddy soils except in soils of Nimaigarh Habigaon (Table 4.8). Humin carbon (C_{Hn}) constituted 11.0-65.1 per cent of soil organic carbon (SOC) in the paddy soils and 30.0-80.0 per cent of soil organic carbon (SOC) in the non-paddy soils (Table 1.2.). The humin C was significantly and positively correlated with soil organic carbon (r= 0.890**), exchangeable Ca⁺⁺ (r= 0.545**) and negatively correlated with soil bulk density (r= - 0.533**) and pH (r= - 0.466*) (Table 1.3.). Humin can be described as the insoluble constituent of soil organic matter (SOM), which remains after extraction of all other

components of SOM which are soluble under acid or alkali conditions (humic and fulvic acids)(R. Swift *et al.*, 2017). Humin makes the largest component of SOM. Therefore, the more soil organic carbon in the soil will mean more accumulation of humin carbon.

The amount of humic acid carbon (C_{HA}) in soil varied from 0.015-0.150 per cent in the paddy soils and 0.015-0.135 per cent in the non-paddy soils (Table 1.1.). Surface horizons of paddy soils contained higher amount of humic acid C (0.120-0.150 per cent) as compared to that of non-paddy soils (0.105-0.135 per cent). Humic acid C constituted 4.3-40.0 per cent of soil organic carbon (SOC) in paddy soils and 8.8-23.7 per cent of soil organic carbon (SOC) in non-paddy soils (Table 1.2.). The humic acid C was significantly and positively correlated with soil organic C ($r = 0.861^{**}$), exchangeable Ca^{++} ($r = 0.444^*$) and negatively correlated with soil bulk density ($r = - 0.644^{**}$) and pH ($r = - 0.453^*$) (Table 1.3.).

The amount of fulvic acid carbon (C_{FA}) in soil varied from 0.030-0.180 per cent in the paddy soils and 0.030-0.300 per cent in the non-paddy soils (Table 1.1.). Surface horizons of paddy soils contained lower amount of fulvic acid carbon (0.135-0.180 per cent) as compared to that of non-paddy soils (0.195-0.300 per cent). Fulvic acid C constituted 33.3-74.0 per cent of soil organic carbon (SOC) in paddy soils and 17.0-50.0 per cent of soil organic carbon (SOC) in non-paddy soils (Table 1.2.). The fulvic acid C was significantly and positively correlated with soil organic C ($r = 0.810^{**}$), exchangeable Ca^{++} ($r = 0.512^{**}$) and negatively correlated with soil bulk density ($r = - 0.540^{**}$) (Table 1.3.).

During rice cultivation, the pH of the surface horizon of paddy soils increases due to submergence. This results in formation of more amounts of humic acid in the surface horizon of paddy soils due to elevated soil pH during rice cultivation. The dominance of fulvic acid carbon over humic acid carbon in the present study is in corroboration with the findings of Borah and Karmakar (1999) and Karmakar and Rao (1999b). The annerobic conditions in the paddy fields and acidic reactions in the studied area are conducive for the formation of fulvic acid over humic acid (Banerjee and Chakravarty, 1977; Karmakar and Rao, 1999b).

Table 1.1. Humus (C_H), humin C_{Hn} , humic acid (C_{HA}) and fulvic acid (C_{FA}) carbon in the soils (%)

Horizon	Depth (cm)	Organic carbon (OC)	Humus (C_H)	Humin (C_{Hn})	Humic acid C (C_{HA})	Fulvic acid C (C_{FA})	C_{HA}/C_{FA}	$(C_{HA}+C_{FA})/C_{TOC}$
<i>P1 (Nimaigarh Habigaon – Paddy soil) : Aquic Dystric Eutrudepts</i>								
Ap	0-15	0.73	0.255	0.475	0.120	0.135	0.889	0.349
Bw1	15-40	0.47	0.195	0.275	0.075	0.120	0.625	0.415
Bw2	40-100	0.15	0.075	0.075	0.015	0.060	0.250	0.500
2Cg	100-120	0.12	0.045	0.075	0.015	0.030	0.500	0.375
<i>NP1 (Nimaigarh Habigaon – Non-paddy soil): Dystric Eutrudepts</i>								
Ap	0-33	0.79	0.300	0.490	0.105	0.195	0.538	0.380
Bw1	33-55	0.63	0.315	0.315	0.090	0.225	0.400	0.500
Bw2	55-80	0.55	0.270	0.280	0.090	0.180	0.500	0.491
Bw3	80-100	0.44	0.120	0.320	0.045	0.075	0.600	0.273
2Cg1	100-190	0.15	0.105	0.045	0.030	0.075	0.400	0.700
2Cg2	190-220	0.09	0.060	0.030	0.015	0.045	0.333	0.667
<i>P2 (Silikha Sanaton – Paddy soil): Dystric Fluventic Eutrudepts</i>								
Ap	0-15	0.54	0.315	0.225	0.135	0.180	0.750	0.583
Bw	15-55	0.42	0.255	0.165	0.105	0.150	0.700	0.607
Bg1	55-90	0.39	0.195	0.195	0.045	0.150	0.300	0.500
Bg2	90-125	0.35	0.180	0.170	0.015	0.165	0.091	0.514
<i>NP2 (Silikha Sanaton – Non-paddy soil): Dystric Eutrudepts</i>								
Ap	0-10	0.57	0.390	0.180	0.135	0.255	0.529	0.684
Bw1	10-35	0.48	0.255	0.225	0.075	0.180	0.417	0.531
Bw2	35-90	0.35	0.135	0.215	0.060	0.075	0.800	0.386
2C1	90-115	0.27	0.075	0.195	0.030	0.045	0.667	0.278
3C2	115-165	0.12	0.060	0.060	0.015	0.045	0.333	0.500
<i>P3 (Khumtai – Paddy soil): Dystric Eutrudepts</i>								
Ap	0-15	0.54	0.330	0.210	0.150	0.180	0.833	0.611
Bw1	15-50	0.42	0.210	0.210	0.075	0.135	0.556	0.500
Bw2	50-100	0.27	0.150	0.120	0.045	0.105	0.429	0.556
Bw3	100-165	0.16	0.100	0.060	0.030	0.070	0.429	0.625
Bg	165-190	0.10	0.089	0.011	0.015	0.074	0.203	0.890
<i>NP3 (Khumtai – Non-paddy soil): Dystric Eutrudepts</i>								
Ap	0-20	0.63	0.435	0.195	0.135	0.300	0.450	0.690
AB	20-27	0.54	0.180	0.360	0.075	0.105	0.714	0.333
Bw1	27-60	0.27	0.150	0.120	0.060	0.090	0.667	0.556
Bw2	60-95	0.17	0.045	0.125	0.015	0.030	0.500	0.265
Bw3	95-125	0.15	0.045	0.120	0.015	0.030	0.500	0.300

The C_{HA}/C_{FA} ratio in the studied soils varied from 0.203 to 0.889 in paddy soils and 0.333 to 0.800 in non-paddy soils (Table 1.1.). The surface soils of paddy soils exhibited higher C_{HA}/C_{FA} ratios (0.750-0.889) as compared to those of non-paddy soils (0.450-0.538) (Table 4.8). The C_{HA}/C_{FA} ratio was positively and significantly correlated with soil organic carbon ($r=0.466^*$), available P_2O_5 ($r=0.384^*$) and available K_2O ($r=0.505^{**}$), and negatively correlated with pH ($r= -0.517^{**}$) and bulk density of soil ($r= -0.375^*$) (Table 1.3.).The C_{HA}/C_{FA} ratio of less than unity indicated dominance of fulvic acid in these soils. The long term rice cropping lends for increased humic acid/ fulvic acid ratio (Zang and He ,2004). The ratio of $(C_{HA}+C_{FA})/C_{TOC}$ varied from 0.375 to 0.890 in paddy soils and 0.265 to 0.690 in non-paddy soils (Table 1.1.). The surface horizons of paddy soils showed lower $(C_{HA}+C_{FA})/C_{TOC}$ ratio (0.349-0.611) as compared to the surface horizons of non-paddy soils (0.380-0.690). These suggest less humification (formation of humus) in the surface horizon of paddy soils. The ratio of $(C_{HA}+C_{FA})/C_{TOC}$ was positively and significantly correlated with available nitrogen ($r=0.426^*$) (Table 1.3.). This suggests that available nitrogen enhances the process of humification. Nitrogen influences and enables the early stage of decomposition because it affects the physiological changes for adaptation of decomposer organisms thus enabling them to function.(Richards, 1987).

Table 1.2. Humus (C_H), humin (C_{Hn}), humic acid (C_{HA}) and fulvic acid (C_{FA}) carbon in soil as percentage of organic carbon

Horizon	Depth (cm)	Humus C (C_H)	Humin C (C_{Hn})	Humic acid C (C_{HA})	Fulvic acid C (C_{FA})
<i>P1 (Nimaigarh Habigaon – Paddy soil) : Aquic Dystric Eutrudepts</i>					
Ap	0-15	16.4	65.1	18.5	34.9
Bw1	15-40	16.0	58.5	25.5	41.5
Bw2	40-100	10.0	50.0	40.0	50.0
2Cg	100-120	12.5	62.5	25.0	37.5
<i>NP1 (Nimaigarh Habigaon – Non-paddy soil): Dystric Eutrudepts</i>					
Ap	0-33	38.0	62.0	13.3	24.7
Bw1	33-55	50.0	50.0	14.3	35.7
Bw2	55-80	49.1	50.9	16.4	32.7
Bw3	80-100	27.3	72.7	10.2	17.0
2Cg1	100-190	70.0	30.0	20.0	50.0
2Cg2	190-220	66.7	33.3	16.7	50.0
<i>P2 (Silikha Sanaton – Paddy soil): Dystric Fluventic Eutrudepts</i>					
Ap	0-15	58.3	41.7	25.0	33.3
Bw	15-55	60.7	39.3	25.0	35.7
Bg1	55-90	50.0	50.0	11.5	38.5
Bg2	90-125	51.4	48.6	4.3	47.1
<i>NP2 (Silikha Sanaton – Non-paddy soil): Dystric Eutrudepts</i>					
Ap	0-10	68.4	31.6	23.7	44.7
Bw1	10-35	53.1	46.9	15.6	37.5
Bw2	35-90	38.6	61.4	17.1	21.4
2C1	90-115	27.8	72.2	11.1	16.7
3C2	115-165	50.0	50.0	12.5	37.5
<i>P3 (Khumtai – Paddy soil): Dystric Eutrudepts</i>					
Ap	0-15	61.1	38.9	27.8	33.3
Bw1	15-50	50.0	50.0	17.9	32.1
Bw2	50-100	55.6	44.4	16.7	38.9
Bw3	100-165	62.5	37.5	18.8	43.8
Bg	165-190	89.0	11.0	15.0	74.0
<i>NP3 (Khumtai – Non-paddy soil): Dystric Eutrudepts</i>					
Ap	0-20	69.0	31.0	21.4	47.6
AB	20-27	33.3	66.7	13.9	19.4
Bw1	27-60	55.6	44.4	22.2	33.3
Bw2	60-95	26.5	73.5	8.8	17.6
Bw3	95-125	30.0	80.0	10.0	20.0

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